

克拉通内走滑断裂成因与控藏机制研究进展

——以塔里木盆地北部为例

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摘要:近年来,在中国多个克拉通盆地腹部识别出成体系发育的走滑断裂系统,这是克拉通盆地内部一种重要的构造样式。针对塔里木盆地北部走滑断裂体系,应用断裂构造解析、断裂发育演化离散元数值模拟、断裂形变有限元数值模拟及断裂核-带结构解剖等技术方法,结合油气井生产动态资料,分析了走滑断裂体系成因与控藏机制,并总结了研究取得的新进展与新认识。研究表明:①塔里木盆地北部走滑断裂体系是盆地中部大型逆冲带向北推挤时形成的伴生调节构造,具有“非共轴挤压、调节区域变形”动力学成因机制;②随着走滑断裂滑移距增大,核部角砾岩类型由裂隙角砾岩、破碎角砾岩演变为杂乱角砾岩和碎裂岩,高阶演化程度的角砾岩可降低断层核部渗透性;③强压扭背景下形成的走滑压脊构造具有“上张下压”纵向应力分布特征,断控储集体主要沿断裂带在深部发育;④走滑断裂相关盐构造耦合-解耦变形特征对油气垂向输导具有重要控制作用;⑤分层变形是小滑移距走滑断裂在深埋条件下的变形特征,可控制油气沿走滑断裂垂向输导后在多个层系分层聚集。

关键词:分层变形;核-带结构;控储、控藏特征;成因机制;克拉通内走滑断裂;塔里木盆地

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Advances in research on the genetic mechanisms of intracratonic strike-slip fault system and their control on hydrocarbon accumulation: A case study of the northern Tarim Basin

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Abstract: In recent years, strike-slip fault system over large areas has been identified in multiple cratonic basins in China, representing a significant structural style in interior cratonic basins. With the strike-slip fault system in the northern Tarim Basin in mind, we apply techniques like seismic-geological analysis of the faulted structure, discrete element numerical simulation of the fault dynamic evolution, finite element numerical simulation of off-fault deformations, and structural analysis of the fault cores and associated damage zones (also referred to as the fault core-damage zone architectures) to study. In combination with well production data, new understandings on the origin of the strike-slip fault system and its control on hydrocarbon accumulation are proposed. The results are as follows. (1) The strike-slip fault system in the northern Tarim Basin is formed as a result of northward thrusting of large thrust belts to accommodate regional shortening under the non-coaxial extrusion in the central Tarim Basin, featuring a dynamic genetic mechanism of non-coaxial extrusion and accommodation of regional deformations. (2) With increasing strike-slip fault displacement, the crackle and mosaic breccias in the fault cores gradually evolve into chaotic breccias and cataclasites. The resulted breccias of high evolutionary degree can reduce the fault core-damage zone permeability. (3) The pressure-ridge structures along the strike-slip faults formed under intense strike-slip transpressional stress, is characterized by extension in the upper part and compression in the lower part, resulting in the fault-controlled reservoirs

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in large scale primarily occurring at depths. (4) The decoupling of the gypsum-salt layer with the overlying strata related to the strike-slip faulting plays an important role in controlling the vertical hydrocarbon migration. (5) Layered deformation, as an inherent characteristic of the strike-slip faults with small displacement at great depths, governs the vertical migration and accumulation of hydrocarbons in multiple layers along the strike-slip faults.

Key words: layered deformation, fault core-damage zone structure, hydrocarbon reservoir- and accumulation-governing characteristics, genetic mechanism, intracratonic strike-slip fault, Tarim Basin

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板块构造理论揭示岩石圈板块相互作用产生的显著变形集中在刚性岩石圈板块边缘,板块内部则相对稳定^[1-3]。大型走滑断裂作为记录刚性岩石圈相对水平运动的典型构造,前期已被广泛关注和研究^[4-6]。这些断裂带发育于板块边缘,在板块汇聚的动力学背景下可延伸至板块内部,因此也被称为板内走滑断裂^[7]。板内走滑断裂将走滑位移从板块边缘传播至板块内部,断裂宽度达数十千米,长度达数百至数千千米,如郯庐断裂、红河断裂和阿尔金断裂等^[8-9]。

不同于板内大型走滑断裂,在克拉通盆地内部也发育千米级滑移距的克拉通内走滑断裂^[10-11]。克拉通内走滑断裂远离活动板块边缘,通常在板内应力集中下形成。由于滑移距小、变形弱、隐蔽性强,前期研究对于克拉通内走滑断裂是否规模发育尚不清楚。近年来,随着盆地内部三维地震资料的不断采集和地震识别技术的持续改进,在中国三大克拉通盆地内部均识别出规模发育的走滑断裂体系^[12]。在塔里木盆地内部,走滑断裂体系多期活动叠加流体改造的控制,在塔中、顺北、托普台和哈拉哈塘等多个地区奥陶系碳酸盐岩中形成了规模储层^[13-16]。除此之外,走滑断裂体系不但垂向沟通深部烃源,而且对油气运聚也具有重要的控制作用^[17-21]。克拉通内走滑断裂发育演化特征与控藏机制也因此成为近年来构造地质学与石油地质学研究的热点^[12-23]。

虽然已发现的走滑断裂断控油气富集区资源量均超过 $10 \times 10^8 \text{ t}$ ^[15, 24],但在走滑断裂体系成因及对油气差异聚集控制机理方面的认识仍存在争议。在成因机制方面,前人研究提出塔里木盆地走滑断裂体系的形成受控于盆地基底结构的非均一性^[25]和原特提斯洋闭合产生的远程挤压应力场^[26],但古生界走滑断裂体系的分布与下伏新元古界裂陷群并无关联性^[27]。在控储机制方面,前期研究揭示了走滑断裂构造破碎成储机制^[28-29],明确了走滑断裂活动形成“核-带”结构,活

动强度大的主干走滑断裂具有多组系“核-带”结构,形成“栅簇状”断控规模储集体^[30-31]。然而,近期勘探实践揭示随着走滑断裂压扭活动强度增大,实钻揭示的断控储集体规模和连通性呈现出新的特点。在控藏机制方面,前期研究探讨了寒武系膏盐岩对走滑断裂垂向输导的影响^[32-33],但在叠合盆地复杂岩性层系垂向叠置发育背景下,走滑断裂分层构造样式对油气运聚及差异富集的控制作用有待进一步深化认识。

本研究以塔里木盆地北部走滑断裂体系为对象,应用地震-地质构造解析、断裂动态发育离散元数值模拟、断裂带变形有限元数值模拟和断裂带核-带结构解剖等多手段、多方法,并结合油气井生产动态资料,提出并总结了走滑断裂体系成因与控藏机制研究新进展和新认识。研究成果不仅为深化塔里木盆地北部走滑断裂成因机制奠定了基础,也为进一步完善断裂构造解析理论、为中国其他海相克拉通盆地走滑断裂体系的研究提供了借鉴,对于明确走滑断控油气富集规律、指导高效勘探开发具有重要实践意义。

1 地质概况

塔里木盆地腹部走滑断裂体系整体具有“大面积展布、多体系并存”的特点(图1)。根据走滑断裂走向和组合样式变化,可将其划分为至少5个体系^[30],包括:①塔中隆起及北坡地区NNE向近平行走滑断裂体系;②阿瓦提坳陷东缘地区NNW向近平行走滑断裂体系;③塔北隆起托普台—哈拉哈塘地区NNE向—NNW向“X”型走滑断裂体系;④塔北隆起于奇地区NE向—近SN向走滑断裂体系;⑤塔北隆起东南斜坡地区NE向—NNW向走滑断裂体系。上述体系在平面上组成了连片展布的走滑断裂系统,覆盖塔北隆起、顺托果勒低隆和塔中隆起等多个构造单元,其南部以吐木休克断裂和塔中南缘断裂等盆内大型逆冲构造带为边界,北

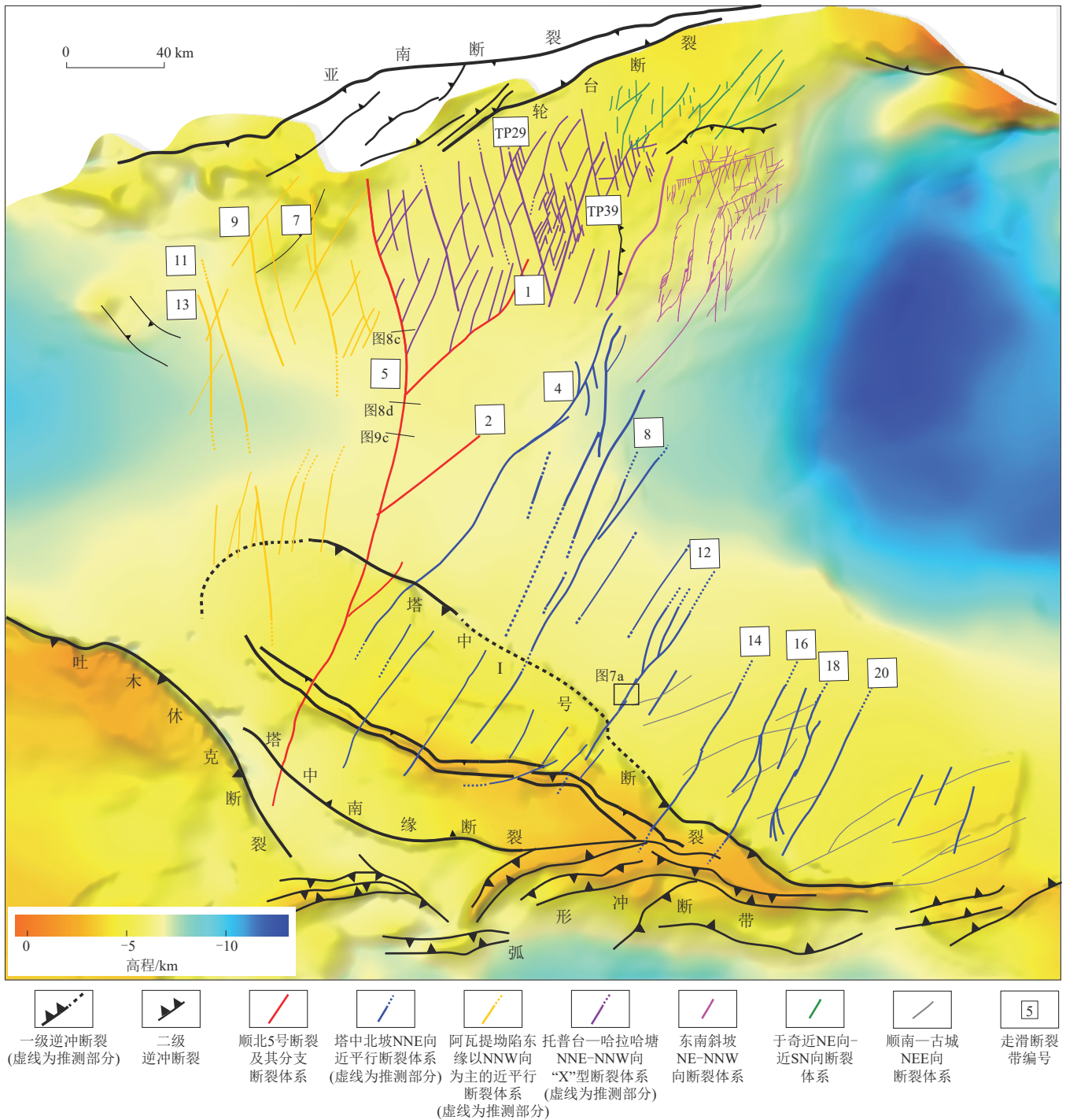


图1 塔里木盆地北部走滑断裂体系分布

Fig. 1 Distribution of the strike-slip fault systems in the northern Tarim Basin

部以雅克拉断凸为边界,东部和西部可见部分断裂延伸至满加尔坳陷和阿瓦提坳陷(图1)。

前期研究表明,上述走滑断裂系统主要形成于中奥陶世末—晚奥陶世末,即加里东中期构造运动^[20, 26]。中奥陶世末,塔里木盆地腹部受与西昆仑造山相关的NE向挤压作用,NW向吐木休克断裂、塔中I号断裂和塔中南缘断裂控制了南部塔中隆起和巴

楚隆起东段的隆升,北部的塔北隆起也开始形成,塔里木盆地广泛发育的稳定碳酸盐台地发生分异,并在古隆起之上形成良里塔格组孤立碳酸盐台地,之后被桑塔木组/却尔却克组淹没台地的复理石沉积覆盖^[34-35]。晚奥陶世末,受塔里木盆地东南缘阿尔金造山和南侧昆仑洋的俯冲,塔中II号断裂和塔中南缘断裂强烈活动,同时形成向北西方向突出的弧形逆冲断

裂系统^[36]。综合分析走滑断裂系统与围限其发育的南部逆冲断裂系统的时-空关联性,推断吐木休克断裂、塔中南缘带和弧形逆冲断裂带等控制了走滑断裂系统的发育。

2 走滑断裂体系成因机制

前期已有学者就走滑断裂与逆冲断裂伴生耦合演化进行了物理模拟研究,揭示了走滑断裂运动学特征差异及其与逆冲断裂的夹角差异会呈现出不同的耦合演化特征^[37-38]。本研究依据塔里木盆地北部走滑断裂体系及围限其发育的盆内逆冲带几何学特征,采用MatDEM软件设置了离散元堆积模型并开展了离散元正演模拟^[39-40],以探索逆冲断裂对走滑断裂体系发育演化的控制作用(图2)。

在模型中,设置了具有不同力学性质的3个区域(图2b)。其中,A区代表顺托果勒低隆,在其东侧发育NNE向近平行走滑断裂体系,西侧发育NNW向近平行走滑断裂体系;B区代表英买力低凸起,发育NNW向近平行走滑断裂体系;C区代表雅克拉断凸,在其南侧沙雅隆起发育NNE向-NNW向“X”型走滑断裂体系。由于B区和C区高强度基底具有上隆特征,因此在模型中被赋予了较高的岩石学强度和黏度系数(表1)。模型的左、右边界分别被施加恒定围

压,与研究区东、西边界实际地质条件相似。模型下边界可分为左下边界和右下边界两部分,分别代表吐木休克断裂和塔中南缘带、弧形逆冲断裂带等走滑体系边界逆冲断裂(图2a, b)。为了模拟与逆冲断裂相似的活动特征,模型的左下和右下边界以梯度位移向区域A推挤(图2b),左下边界和右下边界分两个阶段先后施加梯度位移,这两个阶段与吐木休克断裂和塔中南缘断裂等分别于中奥陶世末和晚奥陶世末活动的相对时限一致。

将正演模拟过程中走滑断裂体系在4个阶段的发育特征与现今展布特征进行对比,结果呈现较好的一致性。在阶段①,当左下边界左端点位移达到 $0.03L$ (L 为模型长度),左下和右下边界过渡点处开始发育弧形断层,断层由南部的走向 20° 过渡为北部的走向 3° 。弧形断层以西发育少量NW向断层,断层走向为 $344^\circ \sim 355^\circ$,弧形断层以东发育1条分支断层,走向为 32° (图3a)。在阶段②,随着左下边界位移增加,早期形成的断层持续生长,当左端点位移达到 $0.10L$ 时,弧形断层贯通研究区发育,且在其东侧可识别出4~5条分支断层(图3b)。与阶段①相比,弧形断层以西的NNW向断层略有逆时针旋转。在阶段③,随着右下边界向上移动,弧形断层以东逐渐形成独立发育的NNE走向断层,走向为 24° (图3c)。在阶段④,随着右下边界位移量增大, 24° 走向断层开始

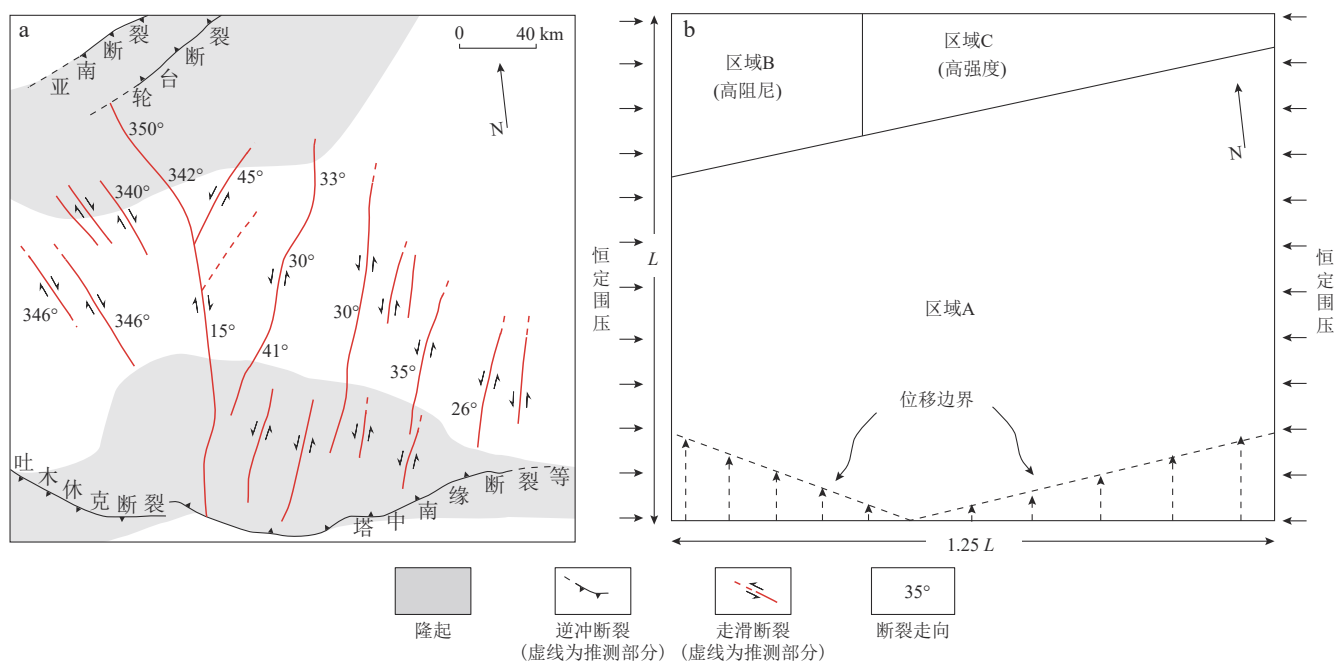


图2 塔里木盆地北部走滑断裂体系分布(a)与离散元模拟边界条件(b)

Fig. 2 Distribution of the strike-slip fault systems in the northern Tarim Basin (a) and boundary conditions for discrete element simulation (b) (L 为模型长度。)

表1 塔里木盆地北部成因机制离散元模拟岩石物理参数

Table 1 Petrophysical parameters for discrete element simulation of the genetic mechanism of the strike-slip fault systems in the northern Tarim Basin

岩石物理参数	区域A	区域B	区域C
杨氏模量/GPa	5.0	5.0	5.0
泊松比	0.2	0.2	0.2
抗拉强度/MPa	2.0	2.0	5.0
抗剪强度/MPa	20.0	20.0	50.0
黏度系数	0.5	1.2	1.2
密度/(g/cm ³)	2.6	2.6	2.6

顺时针旋转,并形成了与之平行的系列NNE向断层(图3d)。

上述模拟结果与现今走滑断裂体系分布的一致性主要体现在3个方面:①弧形断层与顺北5号断裂产状高度相似,且弧形断层以东和以西分别发育NNE向和NNW向断裂,该特征与走滑断裂体系分区发育特征一致;②模拟的NNW向断层(343°~356°)和NNE向断层(33°~35°)与实际的NNW向断裂(340°~346°)和NNE向断裂(26°~41°)走向高度相似;③模拟的弧形断层相关分支断层走向(42°~45°)与顺北1号断裂和2号断裂等分支断裂走向(45°)一致(图3)。

上述模拟结果表明,塔里木盆地北部走滑断裂体系是盆内逆冲断裂向北推挤时形成的伴生调节构造。前人就近平行走滑体系活动造成块体旋转以承载区域挤压缩短开展了深入研究,并辅以断裂体系运动学特征和古地磁学证据进行佐证^[41-42]。鉴于研究区NNE向和NNW向近平行走滑体系必然导致相应块体分别发生顺时针旋转和逆时针旋转,以调节纵向挤压和横向伸展,因此本研究提出了走滑断裂体系的形成受盆内NW向和近EW向两组大型逆冲断裂系分别向东北方向和向北“非共轴”推挤控制,因而具有“非共轴挤压、调节区域变形”成因机制(图4),进一步揭示了塔里木盆地北部走滑断裂是调节区域挤压缩短变形的代表性构造样式。走滑断裂体系在盆地腹部规模发育揭示了区域挤压变形由逆冲断裂系向盆地腹部走滑断裂系递进传播。

在加里东中期走滑断裂体系形成后,不同走滑体系后期差异活动。塔北隆起NNE向-NNW向“X”型走滑断裂体系向南减弱,其中,NNE向体系与塔中隆起NNE向近平行走滑断裂在顺托果勒低隆起南北对接^[42]。在顺北东部地区,主干走滑断裂自西向东滑移距递增,压扭活动增强^[42-44]。

3 走滑断裂断控储集体发育模式新认识

前人对碳酸盐岩走滑断裂内部结构开展了深入研究,揭示了断裂内部变形特征和孔-渗特征的复杂性^[22, 45-46]。随着顺北—富满地区走滑断控油藏勘探实践不断深入,逐步建立了断裂核部破碎角砾带-外侧裂缝带,即“核带”结构发育模式,提出随着断裂活动强度不断增大,断裂内部结构由发育程度较低的裂缝带逐步演化为发育程度较高的破碎角砾带-裂缝带,储集体规模也逐渐增大^[21, 47]。在此基础上,基于深层-超深层测井和取心标定,进一步提出了核-带结构具有栅簇状发育特征,且在拉分段、压隆段等不同分段内部角砾带和裂缝带发育宽度和数量均存在差异^[30-31]。

前期针对顺北地区东部走滑断裂体系运动学特征解析揭示,NNE向走滑断裂体系自西向东滑移距逐渐增大,推测断控储集体规模也随之增大^[11, 43]。然而,随着勘探实践由顺北地区中部主干断裂向东部拓展,获取的岩心资料揭示了顺北8号、12号等断裂发育成熟度更高的核带结构(图5)。前人基于断裂核部角砾带演化过程建立了以角砾大小为主要划分标准的方案^[48],形成以大尺寸角砾(粒径>2.0 mm)、小尺寸碎屑(粒径0.1~2.0 mm)、杂基基质(粒径<0.1 mm)或胶结物含量的三端元分类(图5a),将大尺寸角砾含量大于75%,且角砾呈破碎及弱旋转研磨样式的岩样,定为裂隙角砾岩;将大尺寸角砾含量介于60%~75%,见网状裂纹和角砾破碎、磨圆现象的岩样,定为碎裂角砾岩;将大尺寸角砾含量介于30%~60%,见角砾破碎、磨圆以及无定向分布的岩样,定义为杂乱角砾岩;将大尺寸角砾含量小于30%的岩样命名为碎裂岩,并依据小尺寸碎屑/杂基含量细分为初等、中等以及高等碎裂岩(图5a)。在顺北4号、8号断裂皆可见裂隙角砾岩发育(图5a—c),但在顺北8号断裂可进一步观测到缝网系统发育、角砾呈一定磨圆改造的碎裂角砾岩(图5a, d),以及角砾含量更低(约45%)、杂乱定向分布的杂乱角砾岩(图5a, e),并在局部井段可见大尺寸角砾含量极低(<30%)、杂基/胶结物含量提升并呈一定变形特征的碎裂岩(图5a, f, g),揭示了与顺北4号断裂相比,顺北8号断裂核-带结构演化程度更高。

现今,顺北4号、8号断裂均已建成规模油气产能,其中顺北4号断裂单井单位压降平均油气产量相对高,而顺北8号断裂单井单位压降平均油气产量相对低(图6a)。前人对顺北4号断裂内部结构研究表明,

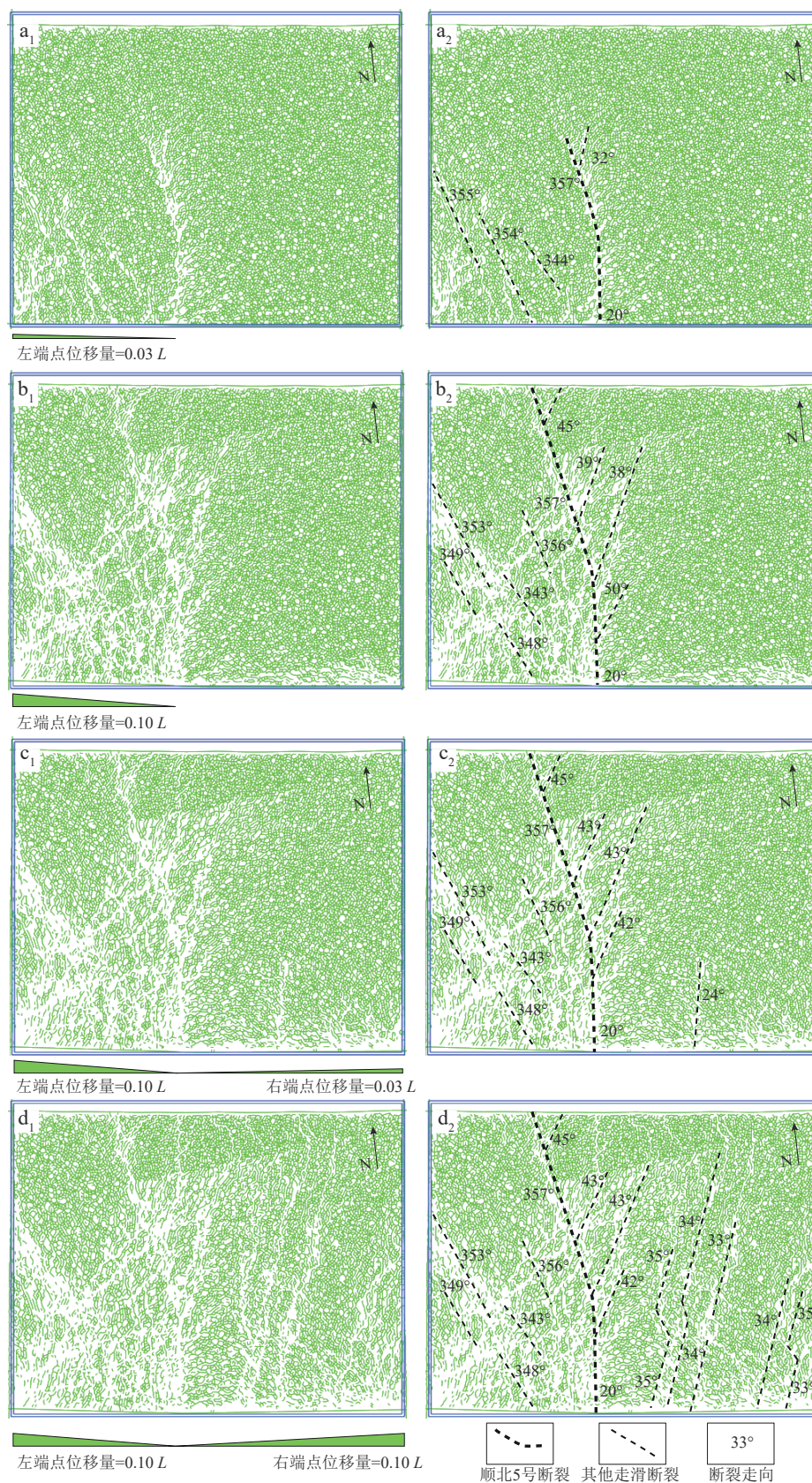


图3 塔里木盆地北部走滑断裂体系成因离散元模拟

Fig. 3 Discrete element simulation of the origin of the strike-slip fault systems in the northern Tarim Basin
 a_1 — d_1 . 分别代表模拟4个阶段解释前的模拟结果; a_2 — d_2 . 分别代表模拟4个阶段解释后的模拟结果
 (L 为模型长度。)

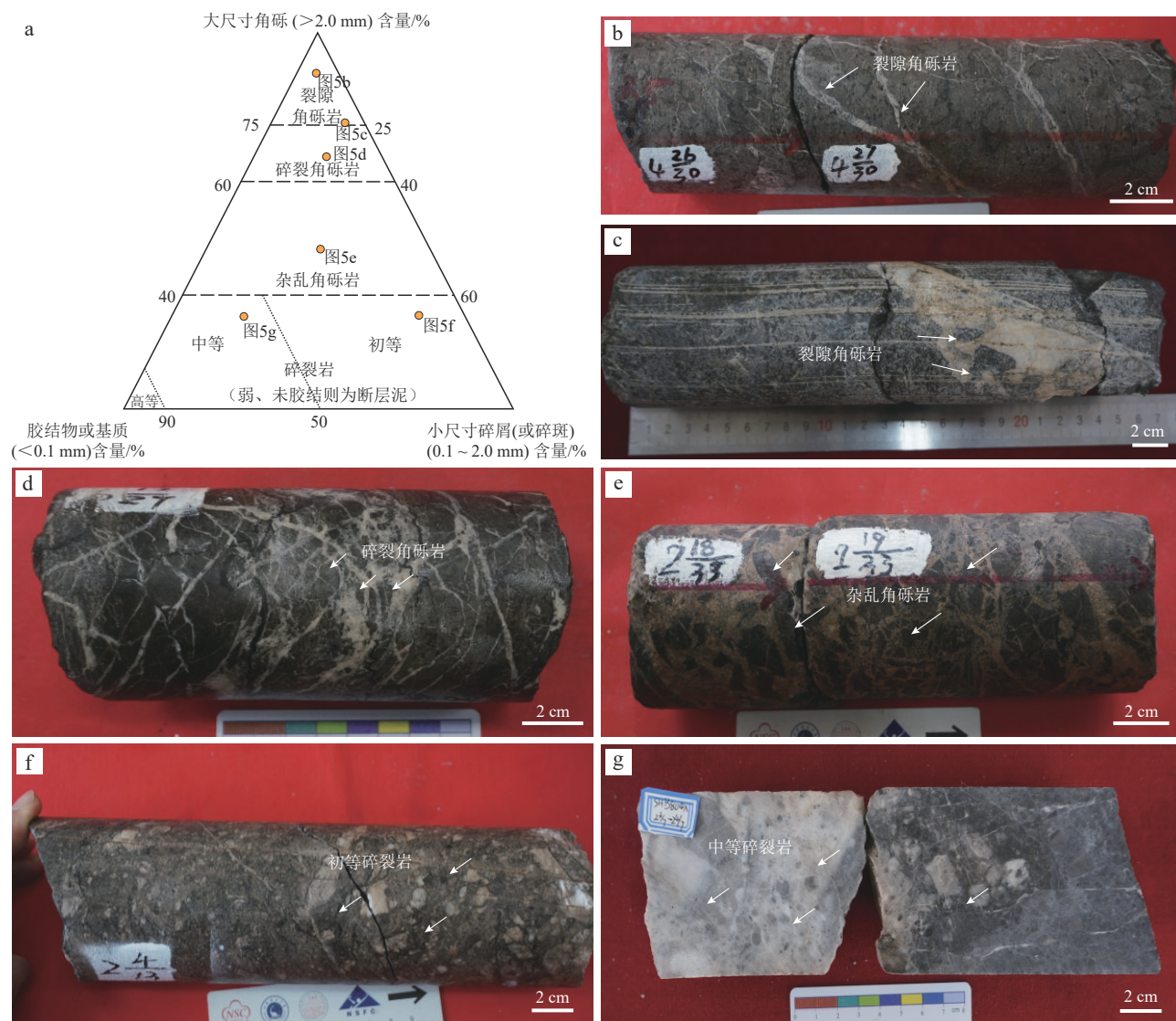


图5 塔里木盆地顺北地区走滑断裂带内部多类角砾岩发育特征

Fig. 5 Developmental characteristics of various breccias within the strike-slip fault zone in the northern Tarim Basin

a. 断裂角砾岩分类三角图, 据文献[48]修改; b. 裂隙角砾岩, 岩心照片, 角砾(85%)、碎屑(5%)、胶结物(10%), SHB82X井, 埋深8 048.50~8 048.70 m; c. 裂隙角砾岩, 岩心照片, 角砾(75%)、碎屑(5%)、胶结物(20%), SHB44X井, 埋深8 180.00~8 180.24 m; d. 碎裂角砾岩, 岩心照片, 角砾(65%)、碎屑(10%)、胶结物(25%), SHB82X, 埋深8 052.00~8 052.21 m; e. 杂乱角砾岩, 岩心照片, 断裂角砾成杂乱分布, 角砾(45%)、碎屑(20%)、胶结物(35%), SHB83X井, 埋深8 324.10~8 324.30 m; f. 初等碎裂岩, 岩心照片, 角砾(20%)、碎屑(60%)、胶结物(20%), SHB84X, 埋深9 088.20~9 088.45 m; g. 中等碎裂岩, 岩心照片, 角砾(15%)、碎屑(45%)、胶结物(40%), SHB804X, 埋深8 554.85~8 555.00 m (图注中括号内数字为含量占比。)

地层左、右两侧平面施加横向挤压应力 60 MPa, 对泥岩地层左、右两侧边界进行 YZ 平面固定约束 (图 7b)。模拟结果揭示, 在泥岩地层中, 地层顶部易产生拉张裂缝, 地层底部易产生挤压剪切裂缝, 在走滑断裂两侧对称分布。在碳酸盐岩地层中, 易产生挤压剪切裂缝, 主要分布于地层底部白云岩及白云岩-灰岩互层中。

模拟结果与真实地震剖面对比揭示, 模型中压脊构造几何形态与实际地震剖面具有相似性 (图 7c, d); 在压脊构造发育过程中, 先存的直立走滑断面在横向

强挤压背景下, 形成压脊构造形态。压脊构造浅部发育张性断面, 应变较小, 而在深部为强挤压变形, 裂缝更为发育 (图 7d)。浅部张性断面易被上奥陶统却尔却克组泥岩充填, 断面深部因两盘强烈挤压作用破碎可形成有利的断控储层。沿断裂走向地震剖面揭示, T_7 界面之上少有“串珠”状强异常反射特征, 异常反射主要分布于深部, 可见模拟结果与该特征一致 (图 7d, e)。以 T_7 界面以下强异常为目标的 W1 和 W2 井均钻遇规模气藏 (图 7d)。综合实钻井地震剖面 and 数值模拟结果分析, 在压脊构造“上张下压”纵

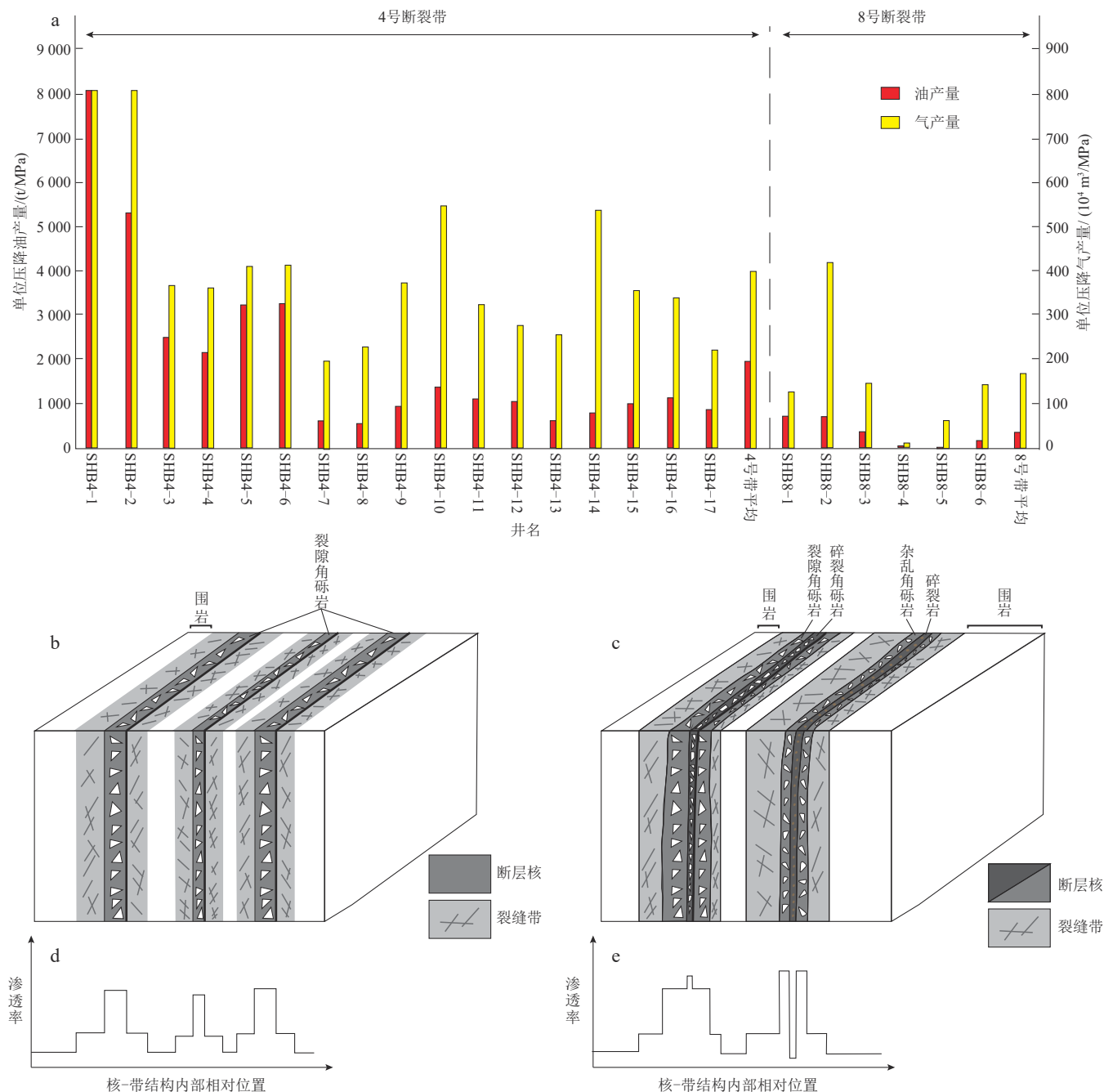


图6 塔里木盆地顺北4号和8号断裂核-带结构单井产能差异与发育模式差异

Fig. 6 Comparison of fault core-damage zone architecture models of Shunbei No. 4 and No. 8 fault zones and their corresponding single-well production capacity per unit pressure drop

a. 顺北4号、8号断裂带产能差异; b, d. 分别为顺北4号断裂核带结构发育模式及核-带结构内部渗透率示意图; c, e. 顺北8号断裂核-带结构发育模式及核-带结构内部渗透率示意图

向应力分布控制下,浅部张性断面易被上覆泥岩充填,在深部挤压应力集中控制下断控储层发育。前期顺南地区主干断裂的钻井资料揭示,奥陶系碳酸盐岩断裂带内部发育裂缝型储层,但胶结充填程度高,同时存在深部酸性热液沿断裂上行溶蚀改造后运移沉淀的特征^[49]。推断顺南地区压脊构造“下溶上填”的成岩改造过程和“上张下压”的变形构造样

式是控制规模储集体在主干断裂深部发育的主控因素。

4 走滑断裂分层变形对油气成藏的控制作用

前人针对塔里木盆地走滑断控油气藏建立了“寒

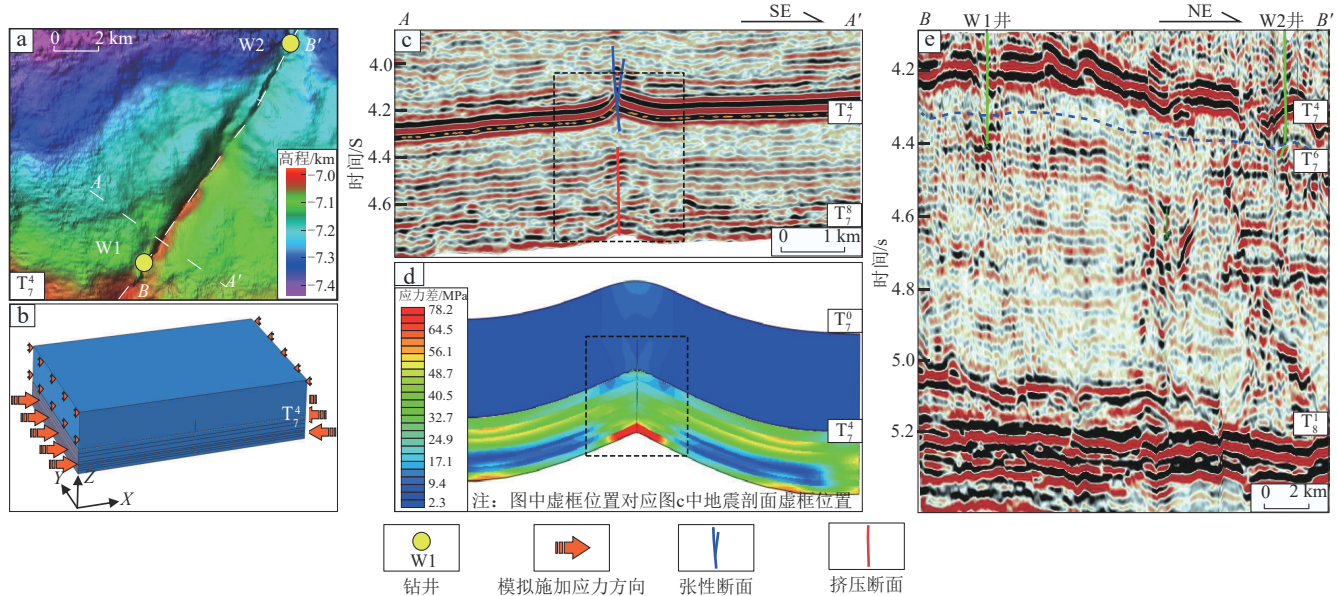


图7 塔里木盆地顺北东部地区主干断裂压脊构造剖面特征与有限元三维模型模拟结果

Fig. 7 Seismic profile characteristics and three-dimensional finite element model simulation results of the pressure-ridge structure along the major faults in the northern Tarim Basin

a. 压脊构造平面特征 (平面位置见图1); b. 压脊构造有限元模拟模型; c. 压脊构造地震剖面特征; d. 压脊构造有限元模拟结果; e. 沿顺北12号断裂走向地震反射特征

武供烃、垂向输导、晚期成藏、断裂控富”成藏模式,指出烃源岩生烃期与断裂活动期的匹配性、断裂穿盐特征等会造成走滑断控油气成藏差异^[15, 28],提出了不同分段样式控制下的“盐加厚”和“盐减薄”变形机制,指出变形可对油气沿断裂垂向输导产生不同程度的影响^[50]。

膏盐岩在高温高压背景下具有从高势区向低势区塑性蠕变的物理特性,在走滑断裂不同分段内部,应力性质差异造成膏盐层厚度变化,膏盐层上、下出现耦合-解耦现象,盐构造变形程度可以用解耦系数进行定量评价(图8a, b)($T=D_1/D_2$; T 为解耦系数; D_1 为变形后的膏盐层厚度; D_2 为变形后的上覆碳酸盐岩地层厚度)^[50]。在压隆段,分段内部压应力集中造成碳酸盐岩顶面挤压隆升,而在深部膏岩层,压应力集中造成膏盐岩逃逸而减薄,变形后的膏盐层和上、下地层具有“X”型特征,解耦系数(T_X)小,输导能力强。反之,在拉分段,分段内部张应力集中造成碳酸盐岩顶面的拉张正断,而在深部膏岩层,张应力集中造成膏盐岩汇聚而加厚,变形后的膏盐层和上、下地层具有“A”型特征,解耦系数(T_A)大,输导能力弱(图8)。实钻揭示,在目的层具有相似“串珠状”强异常背景下,深部具有“X”型变形样式的分段具有更高产能(W3井),而具有“A”型变形样式的分段产能较低(W4井),理论模型与实钻结果一致(图8c, d)。

深层复杂岩性地层中,走滑断裂相关盐构造是典型的分层变形样式。近年来,基于非线性岩石破裂准则与应变模拟,揭示走滑断裂在不同埋藏深度活动时,会使上覆均质体产生不同的形变特征^[51]。在100 m滑移距条件下,走滑断裂在4 km埋深时活动产生的有效应力普遍大于岩石强度,上覆岩体自下而上相对连续变形(图9a)。在8 km深埋条件下,上覆岩体表层和下部有效应力大于岩石强度,而岩体中部有效应力小于岩石强度,断裂上覆地层因而分层变形(图9b)。在塔里木盆地腹部,走滑断裂多期演化普遍处于深埋背景,具有典型的分层变形特征。以顺北5号断裂北段为例,其在加里东中-晚期右行活动,在直立断面之上形成雁列正断层;在海西中-晚期左行继承性活动后,在下伏正断层之上形成又一层雁列正断层,此时,虽然滑移方向出现了反转,但变形并未影响下伏正断层,进而形成了多层不同性质叠置的雁列结构^[52](图9c)。上述分层变形机制导致小滑移距背景下走滑构造在不同层系垂向贯通能力有限。尤其是在大埋深背景下,走滑断裂深部和浅部分层活动,中部趋于稳定,导致晚期生成的高成熟度油气难以向上运移,主要在深层-超深层运聚,因而形成了油气在纵向上分布的有序性(图10)。例如,在同一套通源走滑断裂系统垂向输导控制下,塔深5井在震旦系钻遇油气藏,而于奇14井在白垩系钻遇油藏^[53];在顺北地

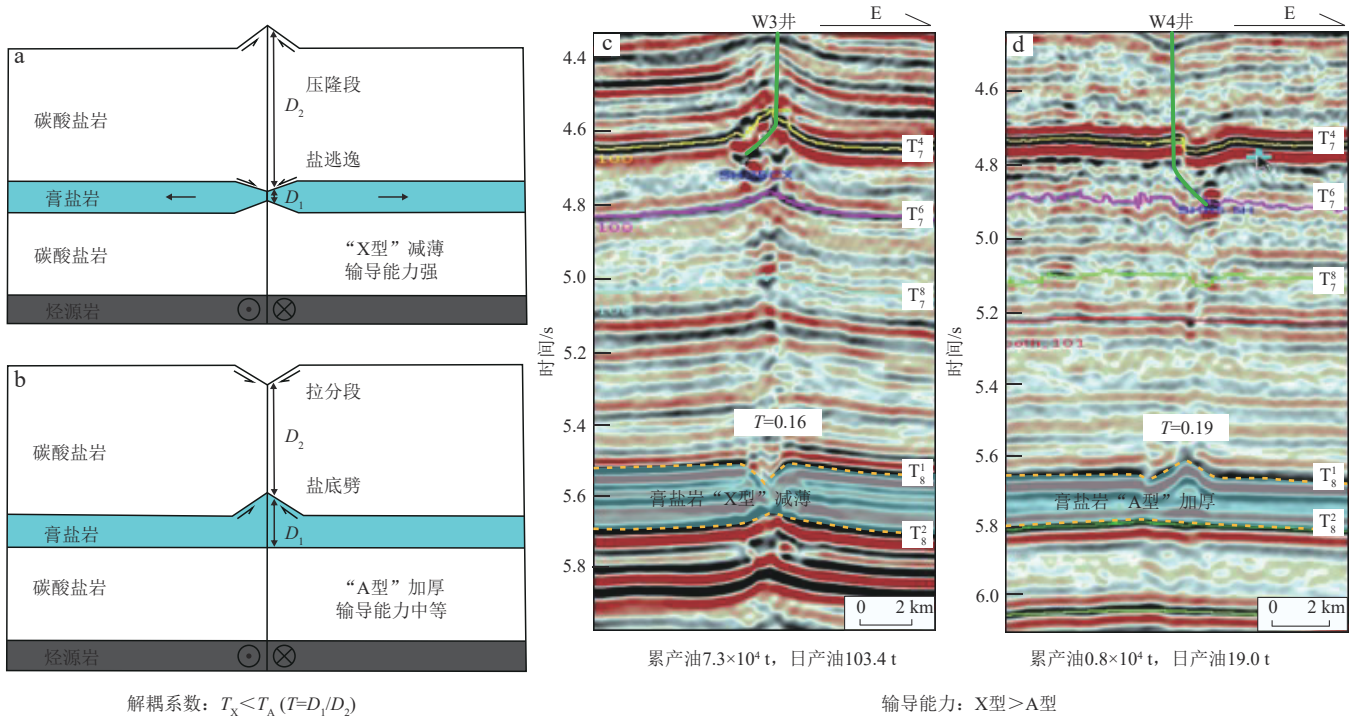


图8 塔里木盆地北部走滑断裂盐构造变形机制与膏盐层厚度变化对油气垂向输导的控制作用(修改自文献[50])

Fig. 8 Deformation mechanism of strike-slip salt structures and the control of gypsum-salt layer thickness on vertical hydrocarbon migration (modified from reference [50])

a. 压隆段盐构造变形机制;b. 拉分段盐构造变形机制;c. 压隆段典型地震剖面;d. 拉分段典型地震剖面

(T 为解耦系数; T_A 为A型加厚解耦系数; T_X 为X型减薄解耦系数; D_1 为变形后的膏盐层厚度; D_2 为变形后的上覆碳酸盐岩地层厚度;剖面位置见图1。)

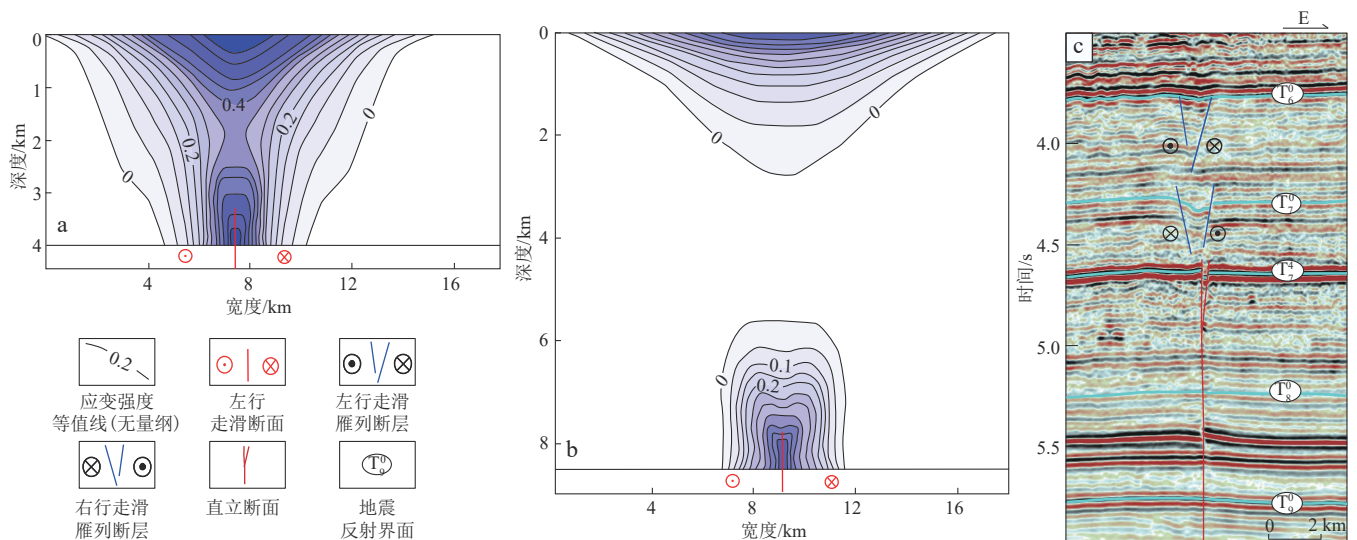


图9 塔里木盆地北部不同深度条件下走滑断裂分层应变分布与分层变形构造样式(a,b修改自文献[52])

Fig. 9 Layered strain distribution and structural styles with layered deformations for strike-slip faults at varying depths in the northern Tarim Basin (Fig. 9a and b modified after reference [52])

a. 浅埋藏条件下走滑断裂应变分布;b. 深埋藏条件下走滑断裂应变分布;c. 过顺北5号断裂北段地震剖面(剖面位置见图1)

区,顺北8号断裂奥陶系钻揭凝析气藏,而在志留系钻揭轻质油藏。走滑断裂早期活动时,断裂通源输导形成顺北8号带志留系油藏,后期深埋后,走滑断裂分层差异变形,中部地层趋于稳定,走滑断裂垂向贯通性

变差,致使晚期凝析气藏在奥陶系碳酸盐岩储层中聚集,形成的走滑断控油气藏得以有序分布。同一套走滑断裂系统控制下形成的多类型油气藏深度高差接近4 000 m。

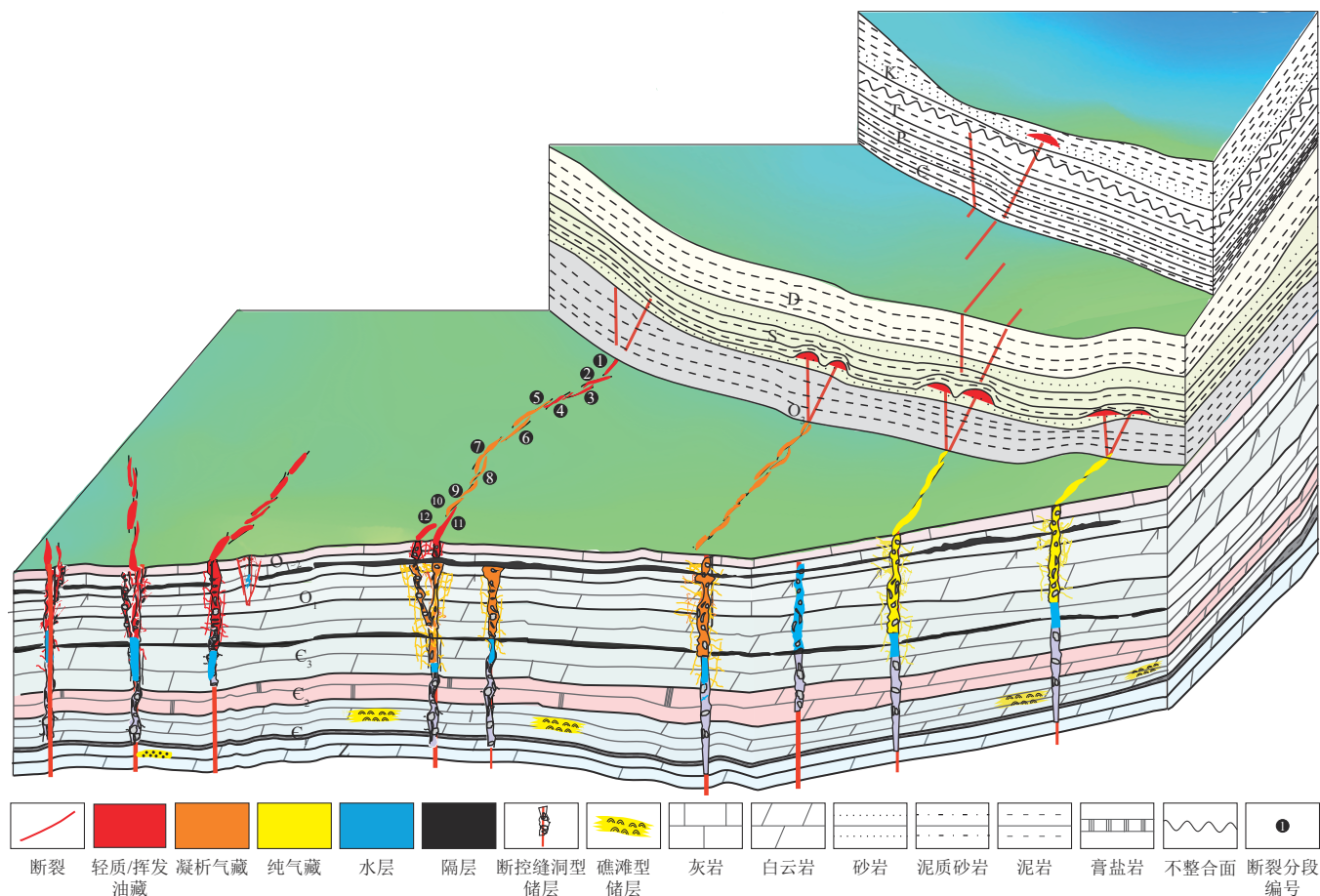


图10 塔里木盆地北部走滑断控油气藏多相态、多层系分布模式

Fig. 10 Multiphase and multilayer distribution patterns of hydrocarbon reservoirs governed by strike-slip faults in the northern Tarim Basin

综上所述,塔里木盆地走滑断控油气藏不仅具有“垂向输导、晚期成藏”的特征,在断裂分层变形控制下,还具有沿断裂分层聚集的特征。断裂通源输导是成藏的第一要素,同一烃源岩系,浅埋藏期,油气沿断裂多层系立体成藏;深埋藏期,受断裂分层变形控制,油气主要在深部近源层系聚集,在塔里木盆地深层-超深层可能发育规模型晚期原生气藏,该认识为克拉通盆地深层-超深层断控油气藏勘探提供了理论依据。

5 结论

1) 塔里木盆内发育NW向和近EW向两组大型逆冲断裂,分别向NE和向北推挤,控制了盆地内走滑断裂体系的发育,是盆内大型冲断带向盆地腹部推挤活动形成的伴生调节构造,具有“非共轴挤压、调节区域变形”动力学成因机制。走滑断裂体系在盆地腹部规模发育揭示了区域挤压变形在空间上由盆缘向盆地腹

部递进传播。

2) 随着走滑断裂滑移距增大,致密碳酸盐岩内部断裂核-带结构向高成熟度演化,角砾岩类型由前期识别的裂隙角砾岩和破碎角砾岩演化为杂乱角砾岩和碎裂岩。高阶演化程度的角砾岩类,特别是碎裂岩会降低断层核部的渗透性,形成相对致密的隔档层,影响断裂内部储集体连通性。

3) 走滑断裂在强压扭背景和复杂岩性地层中均可呈现分层差异变形的特征。强压扭背景下形成的走滑压脊构造具有“上张下压”纵向应力分布特征,浅部张性断面易被上覆泥岩充填,深部挤压应力集中控制断缝储层在深部发育。走滑断裂相关盐构造变形样式受走滑分段性质控制,走滑断裂控制的膏盐层与上、下层系耦合-解耦变形特征对油气垂向输导具有重要控制作用。分层变形是小滑移距走滑断裂在大埋深条件下固有的变形特征;对于同一烃源岩系,浅埋藏期,油气沿断裂多层系立体成藏;深埋藏期,受断裂分层变形控制,油气主要在深部近源层系聚集。

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